Modelling the continuous exchange of nitrogen between microbial decomposers, the organs and symbionts of plants, soil reserves and the atmosphere

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ABSTRACT

Most of the C and N models published over past decades are based on parameters not always linked to the environment and underestimate the role of microorganisms. They are often over-parameterized, which can give multiple solutions for flow calculations between state variables. This work proposes a modelling method centred on the functioning of living organisms in order to calculate flow parameters using data on N stocks in decomposers, plant organs, symbiotic microorganisms, and the soil compartments. The model was settled via a complex N-fixing and intercropping system of durum wheat/faba bean compared to the cropping of pure durum wheat and pure faba bean, all in the context of organic farming invaded by weeds and weeded by hand just before flowering. To avoid perturbation of natural exchanges of C and N, no fertilizer was added from 1997 to 2011.

The equation system defined for the association of any number of plants, as well as parameters previously published for C-flow calculations were used, and only a few parameters specific to N flows were added, and are discussed. The results showed the strong link between N and C in the environment. The equations converge toward an unique set of solutions that is consistent with literature data when available. The labile organic N of microbial origin was modelled as the main potentially available stock. Living microorganisms stored about 1% of total N, which was close to the N stock in faba bean and four times more than stock in durum wheat. Inorganic N was immobilized before flowering in competition with N requirement of durum wheat roots. Net N mineralization, mainly from decomposition of faba bean roots, started too late to improve wheat production. During the cropping period, weeds accounted for losses of 20 kg N ha\(^{-1}\), while the atmospheric N\(_2\) fixation was close to the total microbial immobilization. The model associating microbial and plant flows of C and N in complex plant covers, appears as a robust tool to quantify the exchanges of the earth organisms with the soil and atmosphere. It enables to propose essential recommendations to improve as well agro-ecology as predictions of global changes of C and N stocks.

1. Introduction

Although named Azote by Lavoisier, which means lifeless because inert in gaseous form compared to oxygen, this atmospheric element is probably better named nitrogen (N) in its multiple ionic and organic forms essential to life. It undergoes many transformations of very variable kinetics, inside and outside the soil in the nitrogen cycle (Jetten, 2008), current models of crop production like Ceres (Ritchie and Otter, 1984), Epic (Williams et al., 1989) or Stics (Brisson et al., 1998) integrate management of N conjointly with that of carbon (C) and water. Living plants store a weak part of the global N stock mainly by absorption of mineral N by roots, especially nitrates (Inselsbacher et al., 2013), and fixation of atmospheric N\(_2\) by bacterial biosynthesis of ammonium (Unkovich and Baldock, 2008). Nevertheless, there is much...
less inorganic N available for plant growth than the organic N linked to C forms produced by plants and microbial decomposers (Lin et al., 2000; Pansu and Gauthery, 2006). This underlines the complexity of the N cycle and the need of models describing precisely the N mineralization and immobilization processes linked to evolutions of the microbial biomass and organic substrates. Indeed, the numerous N models reviewed by Manzoni and Porporato (2009) are not always linked to C, sometimes poorly mechanistic and not really centred on the functional ecology of the plant-microbe systems. For Pansu et al. (2009) the published models were over parameterized, two third of them needed parameters not linked to environmental variables, one third did not use explicitly a microbial compartment and none of them considered really the functional role of microorganisms. A new generation was often awaited to connect more deeply the C and N cycles (Gärdenäs et al., 2011), to really express the direct microbial control over decomposition and N exchanges (Blagodatsky et al., 2010; Todd-Brown et al., 2012) possibly at the cellular scale (Gras et al., 2011), if necessary to bridge the gap between linear models and laws of microbial growth (Neill and Gignoux, 2006) and generally to consider more deeply the role of microorganisms in a new green revolution.

Using isotopic data collected from various ecosystems, Pansu et al. (2014) proposed to link the N and C cycles to the functional ecology of soil microbial biomass defined by the MOMOS model (Pansu et al., 2010) parameterized only by 7 kinetic constant all linked to climate, and some of them to soil physical properties and quality of organic inputs. They compared two assumptions (i) microbial homeostasis (constant C:N ratio of soil microorganisms) to reduce the model complexity and the risks of over-parameterization, and (ii) variable microbial C:N ratio to take account of a succession of decomposer communities (Wu et al., 2012) and their diversity (Philippot et al., 2013). The results showed that homeostasis was not always an acceptable hypothesis but could be considered as a valid approximation, especially in hot, well-drained plain areas (Pansu et al., 2014) where prokaryotes are particularly active, confirming the MOMOS expectation to predict the functional ecology of these soil microbial communities (Sikorski, 2015).

Microorganisms representing the largest part of the earth life need plant substrates as C sources with large research questions about the competition or the synergy with plants for the N nutrients (van der Heijden et al., 2008). Corre-Hellou et al. (2009) have modelled crop growth and N accumulation in pea-barley intercrops. Jensen et al. (2015) parameterized only by 7 kinetic constant all linked to climate, and some of them to soil physical properties and quality of organic inputs. They compared two assumptions (i) microbial homeostasis (constant C:N ratio of soil microorganisms) to reduce the model complexity and the risks of over-parameterization, and (ii) variable microbial C:N ratio to take account of a succession of decomposer communities (Wu et al., 2012) and their diversity (Philippot et al., 2013). The results showed that homeostasis was not always an acceptable hypothesis but could be considered as a valid approximation, especially in hot, well-drained plain areas (Pansu et al., 2014) where prokaryotes are particularly active, confirming the MOMOS expectation to predict the functional ecology of these soil microbial communities (Sikorski, 2015).

The field experiment and soil were described in detail by Ibrahim et al. (2016). The soil was a loamy chromic Cambisol with an alkaline pH of 8.2 which greatly differentiate it from the tropical soils of model calibration and validation. This experiment was carried out at the Institut National de la Recherche Agronomique (INRA) Mediterranean station of Mauguio (43°37’12.60”N/3°59’07.12”E), 1 France. Plots of 6 m × 10 m were cultivated with three crop systems (four field replicates) using organic farming methods applied since 1997 without any fertilizer addition (1) durum wheat (Triticum durum Desf. c.v. LA1823) monocrop at a density of 100 ± 23 plants m⁻² (2) faba bean (Vicia faba L. c.v. “Castel”) monocrop at a density of 17 ± 7 plants m⁻², and (3) durum wheat LA1823 - faba bean Castel intercrop at a density of 72 ± 28 durum wheat m⁻² and 16 ± 6 faba bean plants m⁻². The low fertility allowed only these low plant densities but gave optimal conditions to quantify natural exchanges of N between living organisms not masked by external inputs of N. Four whole plants (roots and shoots) of durum wheat and four of faba bean were collected from each plot during growth (1st sampling period) and at maturity (3rd sampling period). A particular attention was necessary to recover the main part of the plant roots using garden forks, several diggings around the selected plants were sometimes necessary and the sampling was not possible on too dry soils. At the same times, two replicate soil samples were collected in 500 mL stainless steel cylinders from the 0–5 cm and 25–30 cm layers from each plot to determine the soil moisture and bulk density.

The soil near the roots was collected from the field, by careful separation by hand from roots, and preserved in iceboxes (4 plots × 3 crop systems × 4 replicates) then gently crushed by hand, without drying, through a 4 × 4 mm grid sieve. The coarse and fine fractions were weighed, recognisable root fragments were separated and joined to their corresponding plant sample. The soil fraction was kept without drying at 4 °C for analysis of microbial N on one part, analysis of total N after drying at 60 °C and grinding at 0.2 μm on the other part. The roots were separated from shoots and washed in deionized water. The root nodules were separated manually and the grains were separated from the shoots. All organs were dried at 60 °C for 2 days and weighed. For subsequent N analysis, samples of each plant organ and soil were grouped and ground to 0.2 mm in a steel planetary ball mill. The microbial N was determined within two days after sampling by fumigation-extraction (Brookes et al., 1985; Vance et al., 1987)

1 There was a mistake in coordinates of Ibrahim et al. (2016) which are not in degrees minutes seconds as indicated but in degrees decimal minutes which corresponds to degrees minutes seconds of this paper.
simultaneously with microbial C (Ibrahim et al., 2016). The fresh thin soil subsample equivalent to 10 g dry soil was fumigated with alcohol-
free chloroform for 18 h. The fumigated sample and a no fumigated control sample were shaken with 30 mL of a 0.5 mol K2SO4 L−1 aqueous solution for 45 min, centrifuged for 10 min and sterilized by filtration using a 0.2 μm membrane syringe. The filtrates were stored in sterile plastic tubes at 4 °C before N analysis which was performed in aqueous phase (Shimadzu TOC-VCSH analyzer). The soil microbial N concentrations (MB-CN) were calculated as the difference between the total N of the extracts of fumigated minus unfumigated divided by a factor kS = 0.54 (Joergensen and Mueller, 1996). N in the unfumigated samples was assimilated to Inorganic-CN.

The root, shoot, nodule, grain organs of each plant as well as a thin fraction of total soil were analysed in elemental analyzer (NA2000, Fisons Instruments) as in Ibrahim et al. (2016). All soil N concentrations (Total-CN in mg g−1, MB-CN, and Inorganic-cN in μg mL−1) were converted to stocks (g N m−2) in the 0–30 cm layer. The total organic N stock in g N m−2 in the 0–30 cm soil layer was:

\[ \text{Total oN stock} = 300 \times bd \times \text{Total-cN(1-Wp)} \times (1-Cf) \]  
\[ (1) \]

where bd was the bulk density of the plot, Wp was the gravimetric water content and Cf was the coarse gravel fraction. For a microbial extract of 30 mL from a subsample of mass mS, the microbial N stock (MB-N) in g m−2 was:

\[ \text{MB-N} = 9 \times bd \times \text{MB-cN} \times (1-Wp)/(1-Cf)/mS \]
\[ (2) \]

and the inorganic N stock was:

\[ \text{inorganicN} = 9 \times bd \times \text{Inorganic-cN} \times (1-Wp)/(1-Cf)/mS. \]
\[ (3) \]

The N stock in each organ of the plants in g m−2 was:

\[ \text{Plant organ N} = m_P-cN \times d \times mP/n \]
\[ (4) \]

where \( m_P \) is the total mass of the plant organ (g), \( n \) is the number of plants sampled, \( m_P-cN \) is the N concentration in the plant organ (g g−1) and \( d \) is the plant density (m−2).

After data collection, the statistical analyses enabled to test variance of sample replicates against variance of the plot replicates, to detect and check external values, to test variance of plot replicates against that of plot modalities and finally to obtain mean values (associated to confidence intervals of Figs. 3–5) for each sampling time and each modality. In a second time, we used these mean values to perform optimization of the parameters of the module of plant growth, in order to find a parameter set which simultaneously predict all these mean data.

2.2. The N transfer equations of the modelling tool

As carbon and nitrogen are closely associated in living organisms, Pansu et al. (2014) proposed to model the MOMOS-N nitrogen cycle in the same way as the MOMOS-C carbon cycle (Pansu et al., 2010). MOMOS (Fig. 1) was defined as a five compartment model (MB, VL, VS, HL and HS) centred on soil microbial biomass (MB) in a matrix equation given by Pansu et al. (2010) for C cycle, by Pansu et al. (2014) for N cycle and submitted to a mathematical analysis by Hammoudi et al. (2014). Each model parameter was linked to temperature \( f(T) \) and moisture \( f(θ) \) functions previously defined and used in Pansu et al. (2010, 2014) and Ibrahim et al. (2016). In this study, the soil water content \( θ \) was predicted using the soil water model SAHEL (Penning de Vries et al., 1989).

In the equation system, the MOMOS-N model matrix \( A_N \) defining N flows was found similar to the MOMOS-C model matrix \( A_C \) defining C flows, only the central term was changed: The central term of \( A_C \) defines microbial respiration when the central term of \( A_N \) defines the microbial exchanges with inorganic N (InorgN, Fig. 1) by a function \( f(\text{Xc,MB}, \text{Xn,MB}) \) of C and N contents of microbial biomass (MB, Fig. 1). Additionally to links between parameters and weather data, the respiration rate \( (k_{res}) \) was found as linked to the soil fine fraction 0–20 μm; the rates of enzymatic digestion of labile \( (k_{vel}) \) and stable \( (k_{vel}) \) plant materials as well as the microbial mortality rate \( (k_{mb}) \) were found linked to the C/N ratio of each organic input (Pansu et al., 2010, 2014). The values in optimum pedoclimatic conditions for the parameters \( k_{VL} \), \( k_{HL} \), \( k_{HS} \), \( k_{SLS} \) remained unchanged from the previous experiments of Pansu et al. (2010, 2014). The present study used again the \( A_N \) matrix and parameters defined for the N cycle (Pansu et al., 2014). Only the initial values of the vector \( x_N \) of the N contents of each compartment (known at the model setup with \(^{15}\text{N} \) tracers), and the continuous inputs of N from the different necromasses (NC) of \( k \) organs of \( j \) plants in the vector \( B_{ij}^k \) were estimated by the general equation:

\[ x_N = f(\text{T}) f(\theta) A_N x_N + \sum_j \sum_k B_{ij}^k \]
\[ (5) \]

where \( A_N \) is the MOMOS-N model matrix (Pansu et al., 2014) and \( x_N = \) is the vector content of each compartment, respectively, \( x_{Ni} \) is the daily output flow of each compartment, and \( B_{ij}^k \) is the vector of external inputs of N. In this unfrizzled intercropping sustained only by restitution of dead materials (organic fertilizer OF N and added inorganic N aN of Fig. 1 = 0 for all \( i \), \( j \) defines the plant (cereal or legume) and \( k \) defines the plant organ or symbiot (litter-NC, root-NC, or nodule-NC). Using \( f_{SS} \) the stable fraction of each plant organ \( j,k \) (data invariant with time which can be estimated by the TAO model from fiber measurements (Thuriès et al., 2002), near infrared (Kaboré et al., 2012), or solid state nuclear magnetic resonance spectrometry (Pansu et al., 2017) and stored in a data base. The \( B_{ij} \) vector of N inputs was adjusted daily to C inputs (Ibrahim et al., 2016) by the balance equation:

\[ B_{ij}^{NF}(t) = \frac{1}{\text{NF}_{ij}^{NC}} - \frac{\text{NF}_{ij}^{NC}}{\text{NF}_{ij}^{NC}} C_{ij} \]
\[ B_{ij}^{S}(t) = \frac{\text{NF}_{ij}^{NC}}{\text{NF}_{ij}^{NC}} C_{ij} \]
\[ B_{ij}^{NF}(t) = B_{ij}^{S}(t) = 0 \]
\[ (6) \]

where \( C_{ij} \) is the daily input of C from organ \( k \) of plant \( j \) (Ibrahim et al., 2016), \( \text{NF}_{ij}^{NC} \) is the CN ratio of a given necromass \( NC \) (assuming data invariant with time which can be estimated by CN analysis of each plant organ and stored in a data base), and \( \text{NF}_{ij}^{NC} \) is the CN ratio of its stable fraction assuming \( \text{NF}_{ij}^{NC} = 50 \) \( \text{NC} \) from NC \( 1^{15} \text{N} \) ratios in the range 200–500 at the end of field incubation of Pansu et al. (2010, 2014). For each incubation time, the flow of the total organic N is the negative of the derivative of the mineralized or immobilized N and is expressed by the sum of the flows \( \dot{x}_{NI} \) of N from each i compartment:

\[ \dot{N} = \sum_{i=1}^{5} \dot{x}_{NI} = -f(\text{Xc,MB}, \text{Xn,MB}) \]
\[ (7) \]

where positive values of the function \( f(\text{Xc,MB}, \text{Xn,MB}) \) correspond to the net mineralization of microbial N and negative values correspond to the net microbial immobilization of inorganic N. The function \( f(\text{Xc,MB}, \text{Xn,MB}) \) of equation (7) was defined using \( \text{NF}_{MB} \), the target value for the C:N ratio of the MB (\( \text{NF}_{MB} \)), assuming a constant \( \text{NF}_{MB} \) ratio throughout incubation (homeostasis):

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2 Sections 2.2 and 2.3 show the N-cycle-specific equations complementary to the previous modelling equations (Pansu et al., 2014; Ibrahim et al., 2016). Readers who are not familiar with the equations can focus only on the results (section 3) and the discussions (section 4).
was the only decomposition parameter adjusted for each plot, all the other parameters and functions linking them to environmental conditions being those of 14C simulations (Pansu et al., 2010) and used in the agro-ecological simulations of this paper. The N initial values of BM, HL, and HS compartments where those found for C (Ibrahim et al., 2016) divided by respective C:N ratios, assuming initial MB C:N ratio = \( \eta_{MB}^{lim} \) and initial HL C:N ratio = \( 3 \eta_{MB}^{lim} \). Only initial HS C:N ratio was optimized (Table 3).

The C flows were previously modelled from the atmosphere to leaves and then to roots and microorganisms, whereas the N flow must be modelled from the soil or atmosphere to roots and symbiotic nodules, and then to the aerial parts (Fig. 1), which induces the equations in 2.3 below.

2.3. Modelling N transfers in plants and microbial symbiosis

Plain bold purple arrows of the flow diagram of Fig. 1 show the N transfers closely associated with most of the C transfers during plant production and microbial decomposition except few specific transfers for C (thin blue arrows) or N (dashed bold purple arrows). Using the previous parameters and relationships of Pansu et al. (2010, 2014) to soil water and temperature data, also to quality of organic inputs for the rates of microbial mortality \( k_{MB}^{resp} \) and microbial assimilation of the labile \( k_{j,k}^{vl} \) and stable \( k_{j,k}^{vs} \) fractions of each organ k of plant j. The respiration rate \( k_{resp}^{j} \) was related to soil texture conjointly with soil water and temperature. The time function \( f_{j}^{1} \) already used for C (Ibrahim et al., 2016) enables to regulate N transfers from shoots to grains, as well as from natural mortality of shoots and nodules of plant j, while the time function \( f_{j}^{2} \) regulates the transfers at harvest. The rates \( \tau_{Am}^{incorp} \), \( \tau_{H} \), \( \tau_{Rm}^{incorp} \) and \( \tau_{mnod}^{incorp} \) used in Ibrahim et al. (2016) for C transfers, regulate also N transfers from natural, harvest, root and nodule mortalities, respectively; \( \tau_{incorp}^{j} \) regulates both C and N transfers from litter to soil by macrofauna and physical disturbance. The TAO model (Transformation of Added Organic materials, Thuriès et al., 2002) split each organic input into labile and stable C fractions linked to stable N fractions by Eq. (6), OF and aiN are possible addition of organic and inorganic fertilizers, respectively (\( \eta_{14C}^{lim} \) = 0 in this study). Grey ellipses and circles show parameters specific to N cycle defined and adjusted in this paper. (For interpretation of the references to colour in this figure legend, the reader is referred to the Web version of this article.)
and = 0 when \( t \neq t_{\text{flowing}} \), \( f^i_j(t, \tau_{i\text{NSR}}) \) regulates inorganic N absorption by roots of each plant at the rate \( \tau_{i\text{NSR}} \) of the root biomass \( y_{i\text{R}} \) expressed in C stock units (Ibrahim et al., 2016). \( \tau_{i\text{NSR}}^{\text{N,mod}} \) regulates the N transfer from symbiotic nodules to root xylem (\( y_{i\text{nod}}^{\text{N}} \) is the nodule N stock = 0 for \( j \neq \text{legume} \)) at rate \( \tau_{i\text{NSR}} \), and the other terms regulate the losses of root N by transfer to shoots at rate \( \tau_{i\text{R}} \) and root mortality at the rate \( \tau_{i\text{Rm}} \) defined by Ibrahim et al. (2016), see Table 1.

The daily change in the inorganic N compartment was the sum of MB mineralization or immobilization (Eq. (7)) and possible inorganic fertilization minus root absorptions and losses from the intercropping system:

\[
\frac{dy_{i\text{N}}}{dt} = f_i(t, x_{i\text{MB}}, x_{i\text{MIN}}, \tau_{i\text{NSR}}) + f_i(t, \tau_{i\text{Rm}}) \sum_j f_j(t, \tau_{j\text{NSR}}) y_j^{i\text{R}} - \tau_{i\text{N,incorp}}^{-1} y_{i\text{N}}
\]

using \( f(x_{i\text{MB}}, x_{i\text{MIN}}, \tau_{i\text{NSR}}) \) defined in Eq. (8), \( f_j(t, \tau_{j\text{NSR}}) y_j^{i\text{R}} \) defined in eq. (9), \( y_{i\text{N,incorp}} \) being the global loss flow of inorganic N in environment, and \( \tau_{i\text{Rm}} \) being the rate of possible addition of inorganic N fertilizer, \( f_i(t, \tau_{i\text{Rm}}, a\text{in}) \) = 0 for all times in this experiment (no fertilization). The N stock in the legume symbiotic nodules was driven by:

\[
\frac{dy_{i\text{N}}^{\text{n,mod}}}{dt} = \tau_{i\text{N,mod}}^{\text{n,mod}} - (\tau_{i\text{NSR}} - \tau_{i\text{Rm}}^{\text{n,mod}}) y_{i\text{N}}^{\text{n,mod}}
\]

where \( \tau_{i\text{NSR}}^{\text{n,mod}} \) is the rate of symbiotic fixation by unit of nodule C stock \( y_{i\text{N}}^{\text{n,mod}} \) (Ibrahim et al., 2016), \( \tau_{i\text{Rm}}^{\text{n,mod}} \) is the nodular mortality and \( \tau_{i\text{NSR}}^{\text{N,mod}} \) the transfer rate to roots (see eq. (9)); all \( \tau \) values = 0 for \( j \neq \text{legume} \).

The N contents \( y_{i\text{N}}^{\text{NSR}} \) of aerial shoots of each plant \( j \) were driven by:

\[
\frac{dy_{i\text{N}}^{\text{A}}}{dt} = y_{i\text{N}}^{\text{A,NSR}} - y_{i\text{N}}^{\text{A}}(f_i(t, \tau_{i\text{N}}^{\text{A}}) + f_j(t, \tau_{j\text{G}})) + f_j(t, \tau_{j\text{G}})
\]

where the first term on the right-hand-side expresses the N inputs from roots, and the second term expresses the N outputs to litter (by natural mortality and harvest) and to grains, assuming that, for a given plant, the same time function \( f^i_j \) drives the C and N transfers to grains and litter; \( f^i_j = 0 \) for all time except at the harvest where \( f^i_j = 1 \) (Ibrahim et al., 2016).

The grain growth takes N from the shoots:

\[
\frac{dy_{i\text{N}}^{\text{G}}}{dt} = y_{i\text{N}}^{\text{G,NSR}}(t, \tau_{i\text{G}}) - f_j(t, y_{i\text{N}}^{\text{G}})
\]

where \( y_{i\text{N}}^{\text{G}} \) is the N exported in grains of each plant \( j \) at harvest time \( t_{i\text{H}} \), and is 0 for all \( t = t_{i\text{H}} \).

The litter \( y_{i\text{N}}^{\text{L}} \) of each plant \( j \) was modelled as the balance of accumulation of aerial parts by mortality and by cutting at harvest, minus the litter transferred by macrofauna to the 0–30 cm soil layer:

\[
\frac{dy_{i\text{N}}^{\text{L}}}{dt} = y_{i\text{N}}^{\text{L,NSR}}(t, \tau_{i\text{L}}) - f_j(t, y_{i\text{N}}^{\text{L}})
\]
gradient descent method.

3. Results

3.1. Modelling measured state variables from N flows

3.1.1. All state variables simultaneously

The plant module enables to predict accurately all the state variables measured at different growth stages in the intercropping system. The fitting by equation (15) (Fig. 2) showed an adjustment significant at $p < 10^{-4}$ of predicted vs measured values and explained 91% of the variance ($R^2$ value on Fig. 2). The values were inside the mean residual sum of square of their distance to the bisector (RSSb on Fig. 2), except:

1. the two last IN points (yellow diamonds in Fig. 2) were slightly underestimated near zero by the model, which was plausible considering the high demand of inorganic N by living organisms;
2. roots (green olive diamond in Fig. 2)) and shoots (green squares in Fig. 2)) of faba bean at 191 day after the seedling; this point was particular in faba bean physiology (Fig. 5); it was the time of grain growth, which took N from shoots and the simultaneous time of depletion of symbiotic nodules which transferred N to roots; the measured root N was exceptionally high at 191 d near the level of the measured shoot N; the modelled increase of root N occurred too late to better fit the measured value, when the modelled decrease of shoot N (transfer to grain N) occurred about 10 days too early to fit the measured values.

Excepting these 4 slight under predictions, the 18 other measured values of state variables were accurately predicted by the above equation system, adjusted to the parameters of Tables 1 and 3 (see section 3.2 below).

3.1.2. Predicting microbial control of organic N

The N stocks and exchanges between organic compartments are
summarized in Fig. 3. The predicted increase of 44 g m$^{-2}$ in total N was not significantly outside the confidence interval of the measured data. It was mainly due to the increase in the labile humus of microbial origin (HL N, Fig. 3a), the model predicted a nearly constant stable compartment (HS N). The amount of total plant debris remaining in the top soil from five sources (roots and litter of durum wheat and faba bean, and symbiotic nodules) was predicted as approximately constant about 0.6–0.7 g N m$^{-2}$ during the entire cycle except for an increase which began around 150 days after sowing, flattening out to 0.9–1.1 g N m$^{-2}$ 200–250 days after sowing, and peaking at 1.4 g N m$^{-2}$ at harvest (orange curve in Fig. 3a).

This plant debris N stock was modelled in Fig. 3b as the incorporation of the daily input from plant organ mortality (See Table 1 remembering the mortality rates of Ibrahim et al., 2016 retained also for N transfers) minus the daily microbial assimilation of debris, which was about 0.1 g N m$^{-2}$ d$^{-1}$ from sowing to 150 d after sowing. After 150 d, the microbial assimilation increased to 0.3–0.5 g N m$^{-2}$ d$^{-1}$, reached a peak of 0.7 g N m$^{-2}$ d$^{-1}$ immediately after harvest and then decreased quasi exponentially to below 0.01 g N m$^{-2}$ d$^{-1}$ one year after sowing (Fig. 3b). The N microbial assimilation rate from plant debris was predicted as being close to the N microbial assimilation rate from labile humus during the 1st step of plant growth, from sowing to 210 d after sowing. The N assimilation rate from labile humus then became increasingly greater than the N assimilation of plant debris until 2 months after harvest. Labile humus was predicted as almost the only source of N for microorganisms 2 months after the harvest (Fig. 3b).

During the entire intercropping cycle, the daily output of N from microorganisms to humus by microbial mortality (Fig. 3b) was predicted as approximately the sum of daily inputs by enzymatic assimilation of plant debris and humus, with variations linked to climate conditions. Microbial inputs (x in Fig. 3c) were significantly correlated with microbial outputs (y in Fig. 3c) using the equation $y = 0.993x$, indicating that daily organic N remaining in microorganisms was only 0.7% of the organic N exchanged by these microorganisms. The model need not an additional arrow and corresponding equation to take microbial assimilation of organic N into account, it simply results from the difference between microbial assimilation of organic materials and the N output by microbial mortality. The other part of N required for microbial functioning was immobilized from inorganic soil N (see 3.1.3 below).

3.1.3. Predicting microbial control of inorganic N

Fig. 4 summarizes the stocks and inorganic N exchanges for microbial decomposers and symbiotic microorganisms. These stocks and exchanges were small compared to the stocks and exchanges of organic forms of N (Fig. 3): microbial N (Fig. 4a) amounted to 0.3–2.8% of the total organic N and 3–23% of N of the labile humus (HL compartment, Fig. 3a), inorganic N (Fig. 4a) ranged from 0.03–3.4% of the total organic N and 0.3–30% of HL N (Fig. 3a). Fig. 4a shows an increase in microbial N which follows the increase in plant restitution during intercropping from 1–2 g N m$^{-2}$ at sowing to 6–7 g N m$^{-2}$ at harvest, and a decrease after harvest when the C supply from plant photosynthesis ceased. The corresponding increase observed and predicted for microbial C by Ibrahim et al. (2016) varied from 10–20 g C m$^{-2}$ at sowing to 60–70 g C m$^{-2}$ at harvest. As for MB-C, the MB-N data were significantly predicted at p < 0.02 (Table 2) by MOMOS with all
predictions within the confidence intervals of the observed values.

The flow of inorganic N (Fig. 4b) was predicted as being net immobilization by microorganisms during the early stages of intercropping until 150 d after sowing when the first periods of net mineralization occurred, followed by swings between mineralization and immobilization, depending on weather conditions, and then by net mineralization after harvest. The cumulated immobilization was predicted as maximal at near 9 g N m$^{-2}$ in the period 150–270 d after sowing, then reducing by mineralization giving a balance of 5 g immobilized N m$^{-2}$ at 360 d (Fig. 4b). The stock of inorganic N (Fig. 4a) was modelled as decreasing significantly under the effect of the microbial and plant uptake from 10 g N m$^{-2}$ at sowing to 0.06 g N m$^{-2}$ 166 d after sowing. It remained low until harvest when it increased again following the decrease in microbial biomass. The predicted values were within the confidence interval of the first measurement and were only slightly under-predicted for the other measurements (Fig. 4a).

The continuous fixation of atmospheric N (Fig. 4c) was predicted as having a quasi linear increase (conjointly with a decrease of inorganic N stock (Fig. 4a)) from 10 to 90 d after sowing reaching a peak of 0.05 g N m$^{-2}$ d$^{-1}$ between 90 and 180 d and then decreasing quasi linearly until about 0.01 g N m$^{-2}$ d$^{-1}$ at harvest when it ceased. Overall, the total N fixation by symbiotic nodular rhizobia was estimated at 9 g N m$^{-2}$ at harvest (Fig. 4c), a value similar to the total immobilization of inorganic N by the other microorganisms (Fig. 4b).

### 3.1.4. Predicting plant uptake and restitution of N

The measured and predicted values of N stored in organs of the intercropping plants are shown in Fig. 5a for durum wheat and in Fig. 5b for faba bean. The N stored in plants was considerably lower than organic N stored in soil compartments and was less than N stored in living microorganisms. The main part of the plant N was stored in the shoots: the shoot N in durum wheat corresponded to 4–24% of microbial N, the shoot N in faba bean corresponded to 2–100% of microbial N. The prediction of grain production of durum wheat and faba bean was significant at $p < 0.01$ (Table 2). The prediction of N in the other plant organs were inside the confidence intervals of the measured values (4 plot replicates $\times$ 4 sampling replicates).

For both durum wheat and faba bean, the N remaining in the roots was about 10% of the N stored in the shoots (Fig. 4a and b). During

### Table 2

Significance of model prediction for each data series (Eq. (16)); see Fig. 1 for variable symbols: MB-N is N of soil microbial biomass, inorgN is soil inorganic N, A$^{FW}$-N is faba bean shoot N, R$^{FB}$-N is faba bean root N, G$^{FW}$-N is faba bean grain N, A$^{FM}$-N is durum wheat shoot N, R$^{MW}$-N is durum wheat root N, G$^{FM}$-N is durum wheat grain N, NS is not significant, ** and *** are significant at $p < 0.02$ and $p < 0.01$, respectively.

<table>
<thead>
<tr>
<th>Variable</th>
<th>$F$</th>
<th>$p$</th>
</tr>
</thead>
<tbody>
<tr>
<td>MB-N</td>
<td>214</td>
<td>**</td>
</tr>
<tr>
<td>inorgN</td>
<td>0.4</td>
<td>NS</td>
</tr>
<tr>
<td>A$^{FB}$-N</td>
<td>3.0</td>
<td>NS</td>
</tr>
<tr>
<td>R$^{FB}$-N</td>
<td>0.9</td>
<td>NS</td>
</tr>
<tr>
<td>G$^{FB}$-N</td>
<td>27.8</td>
<td>**</td>
</tr>
<tr>
<td>A$^{FM}$-N</td>
<td>0.9</td>
<td>NS</td>
</tr>
<tr>
<td>R$^{FM}$-N</td>
<td>0.4</td>
<td>NS</td>
</tr>
<tr>
<td>G$^{FM}$-N</td>
<td>113.9</td>
<td>***</td>
</tr>
</tbody>
</table>

![Fig. 5](image-url)
grain formation N stock was transferred to the grains, which gave a temporary decrease of N in other aerial parts, especially for faba bean (Fig. 5b).

Fig. 5c shows storage of N in plants of about 19 g N m\(^{-2}\) for pure faba bean, 13 g N m\(^{-2}\) for intercropping, and 7 g N m\(^{-2}\) for pure durum wheat, in inverse order of C production (Ibrahim et al., 2016) in these poor fertile plots.

### 3.2. The N flow parameters

All parameters of Table 3 have been optimized together by the Powell’s gradient method for the best prediction of the measured state variables (see 3.1. above). Available inorganic N was continuously regulated by the threshold \(\eta_{\text{lim}}\) for its exchanges with MB when its adsorption by living roots was regulated by the rates \(\delta_{\text{lim}}\) and the time functions \(f_j\) related to physiological requirements of each plant \(j\). The other parameters essential for environmental exchanges where \(\eta_{\text{lim}}\) controlling the amount of N reserve into stable humus, \(\delta_{\text{lim}}\) regulating rate of symbiotic fixation of atmospheric N and \(\eta_{\text{lim}}\) regulating the whole losses of inorganic N. Other parameters of Table 3 regulate the N exchanges between plant organs, from roots to shoots, then from shoots to grains. The other parameters for N exchange remembered in Table 1 had been already optimised for C exchanges.

The maximum absorption of inorganic N was found at 175 d after sowing for durum wheat roots and 225 d after sowing for faba bean roots (Table 3). The cereal N stock was lower than for legume and the N seemed to be transferred more quickly from roots to shoots (Fig. 5a and b). Conversely, the legume shoots contain larger amounts of N than cereals, probably due to the quick transfer from the rhizobia and roots. In the early days of grain formation, the N grain source was shoot N, root absorption of N only reaching a peak in the second stage of grain formation (Fig. 5b). The development of the durum wheat grains occurred before N release by mineralization of dead nodules and roots of the faba bean (Fig. 5b). In these conditions, most of the mineralized N was stored again in the legume organs (Fig. 5c), and did not benefit cereal production.

The N absorption rate for roots was lower for cereal than for legume (Table 3), but must be multiplied by the root biomass or root C to give the total absorption. It might be better to use the specific root surface but Shi et al. (2013) considered that it was better to correlate the root uptake with the root mass, here assimilated to root C in Eq. (9), as the cereal root system is able to explore a much larger soil domain than legume roots (Ibrahim et al., 2016). The rates of N absorption by roots of cereal and legume, were found greater in pure cropping than in intercropping where available N was limited (Fig. 5a), and the sum of root absorption rates of cereal and legume in intercropping was found equal to the root absorption rate of pure cropped legumes.

The N allocation rate to shoots was greater for cereal than for legume. From Table 3, all N daily assimilated by roots of durum wheat was transferred to shoots in intercropping when 1/3 was transferred in pure wheat cropping, illustrating a great N demand of the early flowering cereal in competition with the legume not satisfied by N transfer from faba-bean. In contrast, the N stock and biomass in the shoots were greater for faba bean (Fig. 5b) than for durum wheat (Fig. 5a), probably owing to N fixation, and a lower root absorption rate for cereal than for legume. Nevertheless, the N transfer rate to grains was lower for cereal than for legume (Table 3), again probably owing to a difference between the physiological properties of the cereals and the legumes. In the 1st phase of grain formation, the N transfer from shoots to grains was higher for faba bean (Fig. 5b) than for durum wheat (Fig. 5a), probably owing to N fixation, and a lower root absorption rate for cereal than for legume. Nevertheless, the N transfer rate to grains was lower for cereal than for legume (Table 3), again probably owing to a difference between the physiological properties of the cereals and the legumes. In the 1st phase of grain formation, the N transfer from shoots to grains was higher for faba bean (Fig. 5b) than for durum wheat (Fig. 5a), the peak value of the predicted N in the roots (Fig. 5a) and the N absorbed by roots was predicted as being transferred more quickly for grain growth in cereal than in legume and being stored temporarily in shoots for a shorter time for cereal than for legume.

The fixation rate \(\delta_{\text{lim}}\) of atmospheric N\(_2\) was close to 10 mg N g\(^{-1}\) nod C per day being slightly higher in intercropping than for pure legumes (Table 3), and corresponded to a total fixation of 9 g N m\(^{-2}\) for the entire cycle (Fig. 4c). This experiment showed that the \(\delta_{\text{lim}}\) parameter regulating the N transfer rate from nodules to roots can be removed from the model because the best-fit value was 1 (the maximum value) in both intercropping and pure legume cropping. Each day, all

### Table 3

Specific parameters used to model the N cycle together with the C cycle (Ibrahim et al., 2016).

<table>
<thead>
<tr>
<th>Function</th>
<th>Parameter</th>
<th>Description</th>
<th>Crop</th>
</tr>
</thead>
<tbody>
<tr>
<td>Plant N parameters</td>
<td>(\eta_{\text{lim}})</td>
<td>Root N absorption rate</td>
<td>(g^{-1}) root C d(^{-1})</td>
</tr>
<tr>
<td>(j) (Durum wheat)</td>
<td>(\tau_{\text{NRA}})</td>
<td>N allocation rate to shoots</td>
<td>d(^{-1})</td>
</tr>
<tr>
<td>(\tau_{\text{RRA}})</td>
<td>N transfer rate to grains</td>
<td>d(^{-1})</td>
<td>0.331</td>
</tr>
<tr>
<td>(\tau_{\text{NRF}})</td>
<td>Nodule N fixation rate</td>
<td>g(^{-1}) N d(^{-1})</td>
<td>0.0106</td>
</tr>
<tr>
<td>(\tau_{\text{NRK}})</td>
<td>N transfer rate from nodules to root</td>
<td>d(^{-1})</td>
<td>0.0212</td>
</tr>
<tr>
<td>(\tau_{\text{NNF}})</td>
<td>N transfer rate from nodules to root</td>
<td>d(^{-1})</td>
<td>0.243</td>
</tr>
<tr>
<td>(\tau_{\text{NNR}})</td>
<td>N transfer rate from nodules to root</td>
<td>d(^{-1})</td>
<td>0.728</td>
</tr>
<tr>
<td>(\tau_{\text{NRF}})</td>
<td>N transfer rate from nodules to root</td>
<td>d(^{-1})</td>
<td>0.0078</td>
</tr>
</tbody>
</table>

| N losses | \(\tau_{\text{Nloss}}\) | Rate of losses of inorganic N | d\(^{-1}\) | 0.0024 | 0.0012 | 0.0024 |

| C:N ratios | \(\eta_{\text{lim}}\) | MB C:N threshold ratio for mineralization/immobilization | – | 9.87 | 11.24 | 9.87 |
| \(\eta_{\text{lim}}\) | Initial stable humus C:N ratio | – | 8.38 | 8.14 | 9.11 |

| Time function for N root absorption | \(\delta_{\text{lim}}\) | Optimal time for cereal N root absorption | d | 174.8 |
| \(\delta_{\text{lim}}\) | Deviation time for cereal N root absorption | d | 0.068 |
| \(\delta_{\text{lim}}\) | Optimal time for legume N root absorption | d | 225.1 |
| \(\delta_{\text{lim}}\) | Deviation time for legume N root absorption | d | 0.042 |

\(^a\) Plant \(j\) in pure cropping.  
\(^b\) Plant \(j\) in intercropping.
the N\textsubscript{2} that had been fixed was transferred to the legume roots, then some was transferred to the legume shoots and grains, and some transferred to soil microorganisms by root mortality (perhaps including exudation), which was found to be ten times higher for faba bean than for durum wheat by Ibrahim et al. (2016).

The threshold \(\eta\text{HS}\) of the C:N ratio for microorganisms is the most important parameter for regulating N mineralization or immobilization (Fig. 3b). In this work, in a well-drained system on a plain, microorganism homeostasis was a plausible assumption (Pansu et al., 2014). Net immobilization of inorganic N was predicted when the daily calculated MB C:N ratio was greater than the threshold value of 9.87 in intercropping (Table 3), otherwise, there was net N mineralization. The threshold value was found slightly higher in pure cropping of faba bean than in intercropping, corresponding to higher N mineralization in intercropping. The initial stable humus C:N ratio \(\eta\text{HS}\) was 9.11 (Table 3), which indicates a high reserve of stable organic N, not available at short term for plants, in this soil.

The rate for N losses from the intercropping system \(\tau\text{Nloss}\) was estimated at 0.24% of inorganic N per day (Table 3) which estimates the total losses of 2 g N m\textsuperscript{-2} close to the amount of N measured in the weeds after hand weeding 150 d after sowing. In this system, therefore, the N loss during intercropping was taken to be the N stored in weeds, no supplementary N losses to the atmosphere and groundwater was detected.

4. Discussion

4.1. Robustness and parsimony of the model

All parameter values and relationships previously obtained with \(^{14}\)C isotopic tracers for calibration in two contrasted mountain ecosystems of Venezuela and Bolivia, and validation in 6 very different Venezuelan ecosystems (Pansu et al., 2010) were retained in this study. This study confirms the result of Pansu et al. (2014) based on simultaneous \(^{14}\)C and \(^{15}\)N transfers: the N cycle can be predicted in close association with the C cycle by computing the mineralization or immobilization exchanges between inorganic N and microorganisms. It illustrates the robustness of the model calibrated and validated in various acidic tropical conditions, applied here to calcareous Mediterranean conditions, using the same set of equations and parameters linking environmental variables to microbial functions. Apart from \(\eta\text{HS}\), the N exchange parameter between MB and inorganic N (Eq. (8)), MOMOS does not include parameters unrelated to temperature and moisture, such as the efficiency factors often used in other models. Of published models, it could be the most sensitive to climate change and the most parsimonious (Ockham’s razor, \textit{lex parsimoniae}), in terms of definition of its equations and parameters. The number and diversity of sites where was applied the microbial part of the model regulating the OM transformations in soil authorize to retain the equation system of Pansu et al. (2010, 2014). The parameter \(\eta\text{HS}\) regulates the higher storage of organic N but only in the stable HS compartment which can sustain the life at long term but is not available at short term for crops. Here it is optimized conjointly with the other parameters but measurement methods could be also developed (Pansu and Gautheryou, 2006) for storage in soil databases.

The eco-physiological parameters additionally defined for predicting C transfer between plant organs, microorganisms, the soil and the atmosphere, in this intercropping system (Ibrahim et al., 2016) were also retained for N modelling. But in contrast with the well-tested decomposition model, this plant module is probably improvable, considering more deeply the mechanisms regulating photosynthesis and plant growth, as it is sometimes considered in crop models. Nevertheless, this module enables to predict accurately all the state variables measured at different growth stages in the intercropping system. As transfer parameters optimized by the equation system are often hard to measure experimentally, we propose these equations as a collecting method in various conditions. The eventual improvements will be linked to measurements of new state variables regulating the C and N transfers but will take care to avoid an immediate increase of the number of parameters in order to prevent over parameterization and to preserve the convergence toward unique parameter sets. Decomposers and rhizobia were sufficient for predicting the entire data set of this system, neglecting at this time other possible transfers of N through arbuscular mycorrhiza (Chalk et al., 2006). Equations (9)–(14) are available for all systems of cropping, prairie or fallows with any number of plant \textit{j} growing simultaneously. Nevertheless, they are not tested to describe the aerial and root production of necromass of forest systems, where the state variables of living trees are more difficult to measure but can be estimated by the forest production models.

4.2. C and N flows and yield of the intercropping system

The key to optimizing crop production by using ecological mechanisms will be to adjust the plant N demand linked to C photosynthesis, and microbial production of inorganic N linked to the C availability for microorganisms. In the intercropping example of this study:

1. The amount of C and N in the organic compartments of microbial origin represent the principal reserves in soil, the stable humus HS is the largest reserve which can sustain a latent slow functioning of living organisms; in contrast the labile humus compartment HL is the main stock of potentially available N at short term for cropping’s yields;

2. Living organisms store much less C and N than humus with a higher reserve in microorganisms than in plants. In this study, the decomposer microorganisms immobilized the soil inorganic N in the first 6 months after sowing, then mineralized N when the substrate plant debris increased. The two inorganic N values higher than model predictions can be not only inorganic N: in this unfertilized soil, the K\textsubscript{2}SO\textsubscript{4} extracts can contain a part of organic N forms (anonymous reviewer, personal communication). In addition, wet soil samples could be insufficiently preserved by storage at 4 °C and a slight microbial mineralization could induce a slight overestimation of the measured data explaining the model underestimation. Atmospheric N fixing by symbiotic microorganisms was predicted as being at its peak from 90 to 180 days after sowing. The total fixed N was equivalent to total N immobilization during intercropping; this rate was not measured but introduced as a plausible value before adjusting \(\tau\text{Nfix}\), otherwise the best fit \(\tau\text{Nfix}\) overestimated the N fixation (near 40 g N m\textsuperscript{-2}) compared to most of literature data (Larue and Patterson, 1981). N fixation in Canadian soils was estimated in the range 5–30 g N m\textsuperscript{-2} (50–300 kg N ha\textsuperscript{-1}) depending on the legume species (Yang et al., 2010) excluding faba bean. López-Bellido et al. (2006) estimated the N fixed by faba bean in rotation with cereals in the range 3.9–14.4 g N m\textsuperscript{-2}. Jensen et al. (2010) reported that faba bean had the highest average reliance on \(N_2\) fixation and could save up to 10–20 g N m\textsuperscript{-2} in the amount of N fertilizer required to maximize the yield of crops grown after faba bean. Additionally, it is well documented that phosphorous availability influences the nodule development and N fixation (e.g. Kousas et al., 2005). The experimental plots of this paper were not fertilized for 13 years but we have observed well developed symbiotic nodules during plant preparation, which could explain a relatively high N fixation. For Alkama et al. (2009), the nitrogenase activity can induce local dissolution of P by acidification of root tips favoring P uptake and genesis of new symbiotic nodules on the root hairs. It could be very interesting to extend MOMOS to model the exchanges of inorganic P with the microorganisms, conjointly with the exchanges of N. Further experiments are also required to check whether the above equations of this paper constitute a reliable method for quantifying N fixation. It should be checked against
Ibrahim et al. (2016) showed that the photosynthetic C flows were different for cereals and legumes. According to Trapeznikov et al. (2003), root C of wheat was more developed in poor fertile soils. Cereals mobilized C for root growth to find nutrients with relatively low root mortality (Table 1) but a high loss of C by root respiration (Ibrahim et al., 2016), probably the energy source for growth (Amthor, 2000). The loss of C by root respiration of cereal was much greater than for legume but the transfers of C and N to soil microorganisms by root mortality were largely lower for cereal than for legume (Table 1). The C use efficiency by microorganisms, continuously modelled from all the C inputs minus the C outputs of MB, was found always lower to 0.2 and decreasing (but not significantly) with increasing of the C:N ratio modelled for these inputs as shown for data of Manzoni (2017).

From Ibrahim et al. (2016), intercropping did not increase the total C production in the two plants which was lower than for pure durum wheat in this poor fertile system, contradicting Fujita et al. (1992) but agreeing with the wheat/faba bean intercropping results of Fan et al. (2006). Ibrahim et al. (2016) found optimal times for grain growth of 182 d for durum wheat and 191 d for faba bean, both lower than optimal times of absorption of inorganic N by roots, The durum wheat (Triticum durum) cultivar LA1823 needs most inorganic N just before grain formation at 150–180 d after sowing (Fig. 5a). The developed root system of durum wheat must find inorganic N by soil exploration. At this time the inorganic N was mainly fixed and stored in the legume shoots and grains (Fig. 4b) and latter was transferred to decomposer microorganisms which mineralize inorganic N (Fig. 4b) too late for the cereal requirement.

5. Conclusion

Overall these results show that the main objectives of the introduction section are attained. The general equation system for decomposition previously proposed then validated in various tropical ecosystems appears again valid in alkaline Mediterranean conditions, just after replacement of tropical variables by the new local variables for weather, soil and qualities of inputs from plants. The extension of the decomposition part of the model at regional scale, need only to collect climate data and data on soil texture (fine 0–20μm fraction) or soil pH, and soil water retention, using the optimal values and relationships given in Pansu et al. (2010, 2014). The equation system added to this paper for N transfers through plants and symbiosis appears well complementary of those previously used for C in the same experiment. Though probably improvable in mechanistic terms, it enables to optimize a set of parameter values not depending on initial values. Thus it appears not over parameterized and is now available for use in any agro-ecosystem, using if possible the simplifying hypothesis of microbial homeostasis which is found valid in this work, or the more complex hypothesis of decomposer chain (Pansu et al., 2014). This paper and that of Ibrahim et al. (2016) can be used to give first approximation of transfer rates between plant organs in order to extend the model at regional scales.

Other approximations can be extracted from literature data or from experiments analogous to this work, but with other plant covers. Such combinations of data collection and modelling appear as tools to compare and improve parameters of vegetal production in agro-ecology. Concerning the microbial action on improvement or depletion of the plant growth of this study, the results showed that this vary depending on the phenological stage. During the first six months after seeding, the larger part of inorganic N was immobilized by microorganisms; plants did not take up a great part of inorganic N except atmospheric N₂ for faba bean. After that, the decomposers began to mineralize inorganic N from the dead roots of legume which became available for plant growth, but too late to satisfy the great N requirement of the roots of durum wheat in this unfertilized system. Consequently, research to improve the intercropping of legumes and cereals should select associations of late-flowering cereals and/or early-flowering legumes to improve use efficiency by the cereal roots of inorganic N resulting from decomposition of N rich legume debris. With the plant species of this experiment, the commonly used annual rotation of durum wheat and faba bean (Köpke and Nemecek, 2010; Jensen et al., 2010) must be preferred to intercropping. It should be optimized by complementary research using the modelling protocol to improve C use efficiency by minimizing respiration losses (Manzoni, 2017) and the N use efficiency (Fageria and Baldiari, 2005) by avoiding possible losses of the potentially available N (inorganic N plus living organism N plus labile humus N) between the legume and the cereal crops. The modelling protocol of this paper jointly to that of Ibrahim et al. (2016) appear as a new tool to study the C and N transfers and their relations in all cultivated and natural systems. Here, it was applied to two plants in intercropping in comparison with pure cropping, but the equations authorize transpositions to other systems with any number j of associated plant species. The modelling method enables the simultaneous calculation of exchange parameters of great importance in plant physiology, by optimization against state variables of C and N stocks much more easy to collect than flow data. For example, it should be very interesting to extend the above study to a large panel of experiments comparing modelling results of different agricultural choices. If possible, to compare with other measurement methods of N and C flows (as isotope methods) in order to collect essential parameters such as growth rates of plants, allocation rates to plant organs, adsorption rates by roots, fixation rates by symbioses, or rates of losses for crops. And to attempt extending the studies to the modelling of the exchanges of phosphorous between the living organisms and the soil compartments.

This mechanistic modelling assigns to microbial decomposers their key role in linking the growth of the different plant species and symbioses in interaction. An experimental or bibliographic collecting of parameters defined in this paper conjointly with those of Ibrahim et al. (2016) and Pansu et al. (2010, 2014), see compilation in supplementary material annexed to this paper, will enable to set out a strong data base for controlling the C and N exchanges of the life at regional scale. Already recognised as generalizable for decomposition; the MOMOS model will be an invaluable tool of prediction of C and N transfers between microbe and plant organisms, which will greatly improve predictions of the reserves linked to the earth life and acting on climate change.

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Appendix A. Supplementary data

Supplementary data related to this article can be found at https://doi.org/10.1016/j.soilbio.2018.06.011.